

Methane optical density measurements with an integrated path differential absorption lidar from an airborne platform

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Abstract. We report on an airborne demonstration of atmospheric methane (CH₄) measurements with an Integrated Path Differential Absorption (IPDA) lidar using an optical parametric amplifier (OPA) and optical parametric oscillator (OPO) laser transmitter and sensitive avalanche photodiode detector. The lidar measures the atmospheric CH₄ absorption at multiple, discrete wavelengths near 1650.96 nm. The instrument was deployed in the fall of 2015, aboard NASA's DC-8 airborne laboratory along with an in-situ spectrometer and measured CH₄ over a wide range of surfaces and atmospheric conditions from altitudes of 2 km to 13 km. In this paper, we will show the results from our flights, compare the performance of the two laser transmitters, and identify areas of improvement for the lidar.

Keywords: lidar, spectroscopy, methane, optical parametric amplifiers, optical parametric oscillators, airborne instruments.

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1 Introduction

Methane (CH₄) is the second most important anthropogenic greenhouse gas (GHG) with a higher radiative forcing potential than Carbon Dioxide (CO₂) on a per molecule basis¹, making anthropogenic CH₄ a critical target for mitigation. The current CH₄ global mixing ratio is 1852 parts per billion (ppb)^{2, 3}. Anthropogenic CH₄ is responsible for a significant portion of the global warming produced by all well-mixed greenhouse gases and contributes to the formation of ozone⁴, another GHG and air pollutant.

Despite the critical importance of CH₄ for climate, the existing CH₄ observing network has proven inadequate to constrain global, regional, and sectoral sources, and explain observed trends and variation in atmospheric CH₄ over the last few decades. Therefore, there is a critical need for CH₄ observations for constraining the strength and distribution of methane's sources, including natural (e.g., wetlands) and anthropogenic (e.g., energy sector) ones. For instance,

much of the year-to-year variations in methane's global growth rate are likely from variations in wetland emissions and part of the recent increasing trend in methane's growth rate may be associated with increased energy extraction activities^{5, 6}. An adequate CH₄ observing network is necessary to monitor the interaction between the carbon cycle and climate change, such as the potential release of CH₄ from stored carbon reservoirs (e.g., Arctic and boreal soils) and changes in natural emissions. The current CH₄ observing network does not provide the necessary data to understand and constrain methane's sources, such as from permafrost thaw, wetlands, which challenges our ability to make confident projections of future climate. The importance of measuring CH₄ is also reflected in the last National Research Council Decadal Survey for Earth Science⁷ and the recent report by the Carbon Climate Workshop⁸.

Our current understanding of CH₄ distributions and processes is founded mostly on precise and accurate ground-based, in-situ measurements from global monitoring networks^{9, 10}. The location and frequency of these measurements is, however, very sparse on a global scale and is even sparser at high latitudes where the thawing Arctic permafrost is of particular concern. Large quantities of organic carbon are stored in the Arctic permafrost and a warming climate can induce drastic changes in carbon emissions and a subsequent positive feedback mechanism that can significantly accelerate climate change¹¹.

Global measurements from satellites are available from passive optical sensors AIRS¹², SCIAMACHY^{13, 14}, TES¹⁵, IASI¹⁶, and GOSAT¹⁷, but currently lack the required sensitivity to derive regional CH₄ sources. Passive sensors measuring reflected sunlight are limited to sunlit areas of the planet and their sensitivity falls off at low sun angles, increasing cloud cover, aerosol scattering, and low surface reflectivity. Recent observations indicate that the thawing Arctic

permafrost is active even during the cold season¹⁸ highlighting the need for continuous sampling at high latitudes even in the winter months.

The benefit of active sensing missions is that they provide global CH₄ measurements where they are really needed: in the absence of sunlight (i.e., at night and at high latitudes in all seasons), in the presence of scattered or optically thin clouds and aerosols, over land and water surfaces, and with higher accuracy and precision than currently available. Active measurements using laser remote sensing technology will be a key step in obtaining measurements of CH₄ from orbit with sufficient coverage, sampling, accuracy and precision to address key science questions. The French Centre National d'Etudes Spatiales (CNES) in collaboration with the German Aerospace Centre (DLR) are developing an active methane mission called MERLIN (Methane Remote Sensing Lidar Mission) scheduled for launch in 2021^{19, 20}. The MERLIN mission targets an 8-36 ppbv relative random error in the methane column abundance with a 50 km horizontal resolution. At NASA Goddard Space Flight Center (GSFC), we have been developing an active, airborne lidar to measure atmospheric methane using Integrated Path Differential Absorption (IPDA) as a precursor to a space mission to measure CH₄ from orbit.

2 Instrument Description

An IPDA lidar measures the absorption of laser pulses by a trace gas when tuned to a wavelength coincident with an absorption line²¹⁻³¹. Using the instrument in a sounding (surface reflection) mode which enables integrated column trace gas measurements from orbit with relatively modest laser power.

The GSFC IPDA lidar uses a tunable, narrow-linewidth light source and a photon-sensitive detector coincident with a CH₄ absorption at 1650.96 nm. The CH₄ spectrum at 1650.96 nm is well suited for active remote sensing. The CH₄ line is mostly isolated from adjacent CO₂ lines

and there is very little water (H₂O) vapor interference. The MELRIN line at 1645.55 nm is less suitable for our technique because it is interfered with by H₂O vapor at ~1645.47 nm and it is wider than our line (~56 pm vs. ~36 pm), an important consideration because it increases the tuning requirement for our laser transmitter. Fig. 1 shows the two-way atmospheric transmittance spectrum around 1650.96 nm from a 400 km orbit using the 2008 HITRAN database³² and a US standard atmosphere.

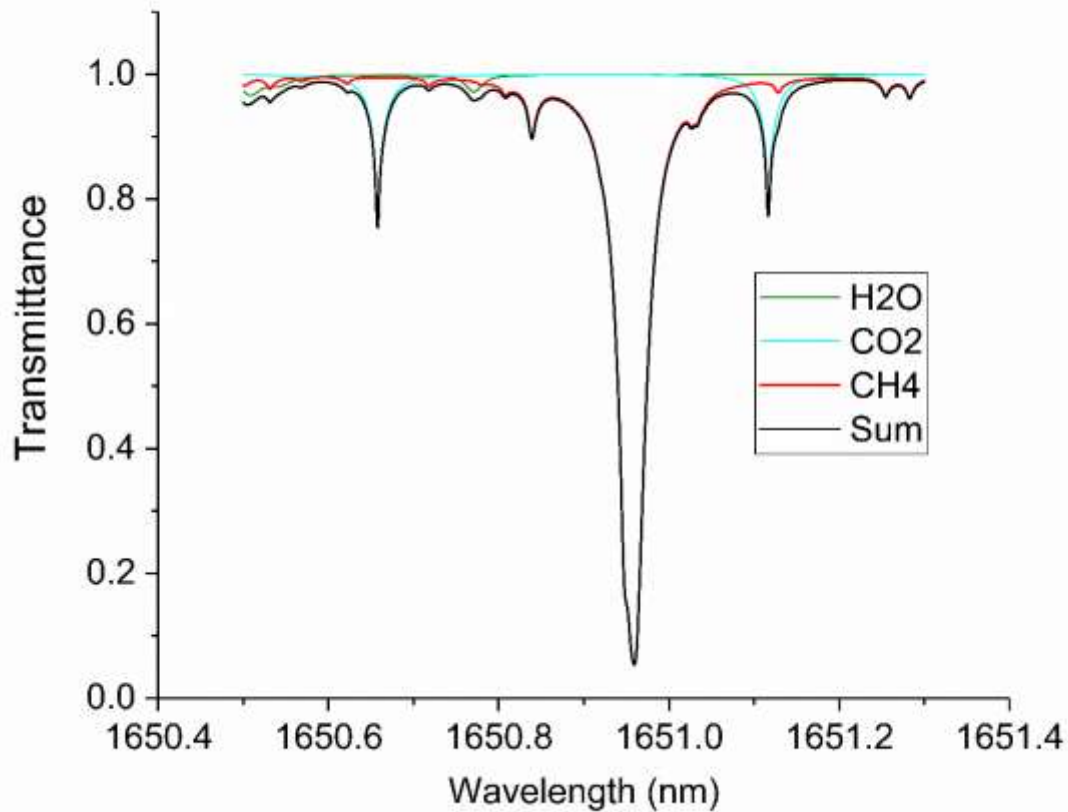


Fig. 1. Two-way atmospheric transmittance near 1650.96 nm a from a 400 km orbit using the 2008 HITRAN database and a US standard atmosphere. The CH₄ line is mostly isolated from adjacent CO₂ and H₂O vapor lines.

Although in principle, only two wavelengths (“on” and “off” the line) are needed to determine the transmittance through the atmospheric column, our technique uses multiple wavelengths to

probe the absorption feature. Using multiple wavelengths can reduce errors that may affect the measurement precision³³, measure the spectral shift of the line with changing atmospheric pressure³⁴, generate atmospheric backscatter profiles of the entire column³⁵, and enable retrievals of trace gas mixing ratios above and below the planetary boundary layer³⁶.

An early version of our instrument²⁵ flew in 2011. The major differences between the system in ref. 25 and the new instrument are: 1) the detector: in 2011 we used a very low (<1%) quantum efficiency (QE) photomultiplier tube (PMT) with very limited dynamic range. The new instrument used an enhanced avalanche photodiode (e-ADP) with ~90% QE. 2) The type and energy of the transmitter(s): In 2011, we used a low energy Optical Parametric Amplifier (OPA) laser transmitter with pulse energy of ~10 μ J. The new airborne lidar used an OPA and an Optical Parametric Oscillator (OPO) with pulse energies of ~25 μ J and ~250 μ J respectively. 3) The opto-mechanical layout and data acquisition system were completely redesigned and considerably improved. As a result, spurious effects such as etalon fringes were dramatically reduced which improved the precision of the instrument.

Our new airborne IPDA lidar used two different laser transmitters. The first is an Optical Parametric Amplifier (OPA) and the second is an Optical Parametric Oscillator (OPO). Only one laser transmitter is used at a time by using a movable mirror to select the desired transmitter. A simplified block diagram of our lidar is shown in Fig. 2 and is based on our previous work with optical parametric generation³⁷.

The OPA, used 20 wavelengths, but was simpler to implement than the OPO, because it did not require an optical resonator cavity, was easier to align and tune, and used only two seed lasers. However, it is extremely difficult to scale the OPA energy to the level needed for space and still maintain a narrow linewidth. Depending on the receiver size and other instrument parameters

115 we calculate that approximately 600 μJ is needed to obtain a measurement with a 0.5% precision.
116 The underlying reason for the wider linewidth is the large mismatch between the seed and pump
117 energies, which makes it very difficult to amplify the seed with the desired spectral
118 characteristics. If the seed laser power can be significantly scaled up and back-conversion can
119 be suppressed then it may be possible to achieve energies of 600 μJ out of the OPA with the
120 desired spectral characteristics.

121 In the OPO, narrow linewidth is achieved by using an optical resonator cavity, which also
122 enhances the energy of the non-linear conversion. Our OPO uses five wavelengths and a 1.2 mJ
123 GSFC-built solid-state pump laser with a triangular optical ring cavity. The OPO energy could
124 be scaled to space ($\sim 600 \mu\text{J}$) and maintain a narrow linewidth with a suitable higher energy pump
125 laser and improved optical design. However, the OPO currently requires a separate seed laser
126 and complex optical phase-lock loops for each wavelength used.

130 Oscillator (OPO). The transmitters use DFB diode lasers for seed lasers but different pump lasers. Only one laser
131 transmitter is operating during flight.
132

133 The first transmitter option (OPA) consists of a magnesium oxide-doped periodically poled
134 Lithium Niobate (MgO:PPLN) crystal which is pumped by a pulsed single-frequency 1064 nm
135 laser and seeded by a continuous-wave (CW) 1650.96 nm laser diode. The pump laser is a
136 custom burst-mode Yb fiber laser from Fibertek Inc., based on a Master Oscillator Power
137 Amplifier (MOPA) configuration³⁸. The pump laser was optimized for high peak power and
138 generated 600 μJ in a burst pulse. Each burst pulse consists of twenty individual 3 ns pulses
139 separated by 85 ns with the individual pulse energies in the burst varying from 2 to 10 μJ . An
140 example of the OPA burst pulse is shown in Fig. 3.

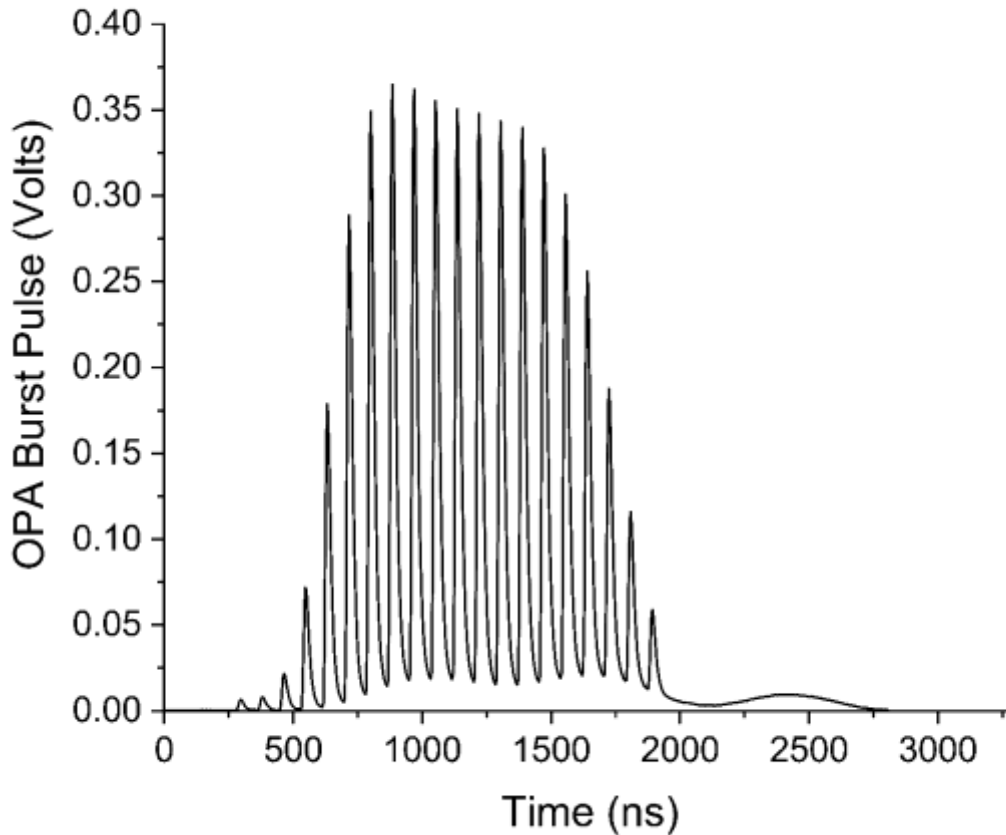


Fig. 3. Example of temporal shape of the OPA burst pulse from the energy monitor detector showing individual pulses within the burst pulse.

The OPA output varies non-linearly with the peak power of the pump so the variation in the individual pump pulses resulted in very low conversion in the OPA of the low energy pulses. The linewidth of the OPA was ~ 500 MHz. The pump laser was delivered with a bare, large mode area (LMA) fiber output that optimized the power output but was not suitable for flight. Prior to our flights, we connectorized the output and the burst pulse energy was reduced to 350 μJ per burst pulse.

Two distributed-feedback (DFB) CW diode lasers, a master reference and a scanning seed, from NEL America (NLK1U5FAAA), are used in the OPA. The wavelength of the master reference

153 laser is locked on the absorption peak at ~ 1650.96 nm using a 16.5 cm cell containing ~ 40 mbar
154 of CH_4 . The locking technique is the same for both the OPA and OPO, and is described by
155 Numata^{39, 40}. It is based on the technique used by Pound–Drever–Hall⁴¹ and is similar to the
156 technique used by Fix⁴². We estimate that the long-term drift of the master laser frequency is \sim
157 2 MHz, based on our experience with similar DFB seed lasers for CO_2 .
158 The scanning seed laser is tuned over the CH_4 line by rapidly scanning the laser current. The
159 beat signal between the master reference laser and the scanning seed laser is measured by a
160 frequency counter and recorded by the data acquisition system. The frequency of the beat signal
161 is converted into the OPA wavelengths in post-processing.
162 The scanning seed and pump laser beams are combined with a beam combiner and focused
163 through the PPLN crystal. The temperature of the PPLN crystal can be temperature-tuned to
164 optimize the phase matching at the seed wavelength. The unconverted pump beam at 1064 nm is
165 separated from the signal beam at 1650.96 nm using a dichroic mirror and directed into a beam
166 dump. A small part of the OPA beam at 1650.96 nm is also directed through an 8 cm reference
167 cell containing ~ 170 mbar of CH_4 for calibration purposes and a blocking filter prevents any
168 remaining 1064 nm radiation from exiting the aircraft. The main OPA output beam is directed
169 through a beam expander to reduce its divergence. The final output energy of the OPA
170 transmitter exiting the aircraft was approximately 25 μJ per burst pulse and 20 wavelengths were
171 used in each wavelength scan to sample the CH_4 lineshape.
172 The second transmitter (OPO) consists of another temperature controlled PPLN crystal inside a
173 three-mirror cavity. The temperature of the PPLN crystal can be temperature-tuned to optimize
174 the phase matching at the seed wavelength. The OPO is pumped by a pulsed single-frequency
175 1064 nm Nd:YAG laser and seeded by five CW DFB lasers at ~ 1650.96 nm.

176 The pump laser of the OPO is a custom-made GSFC single frequency Nd:YAG laser with a ~60
177 ns pulse width and maximum energy of 1.2 mJ per pulse at a 5 kHz repetition rate.

178 The same master reference seed laser used in the OPA is used for the OPO. Part of the phase-
179 modulated master reference beam (labelled “locking beam” in Fig. 2) is used to lock the OPO
180 cavity using a cavity length-control servo and a piezo electric transducer (PZT). Four additional
181 slave DFB diode lasers are offset-locked to the master reference laser by an integral number of
182 the OPO cavity free-spectral range (FSR) using the beat signal from four detectors and four
183 optical phase locked loop servos and a Rubidium frequency reference. Thus, the OPO samples
184 the CH₄ absorption at five wavelengths (one master and four slave). The wavelengths are
185 selected by switching fast (NanoSpeed™) 1x2 fiber optic switches made by Agiltron.

186 After exiting the OPO cavity the unconverted pump beam at 1064 nm is separated from the
187 signal beam at 1650.96 nm using a dichroic mirror and directed into a beam dump. After the
188 dichroic mirror, a small part of the OPA beam at 1650.96 nm is directed through a 5 cm CH₄
189 reference cell containing ~260 mbar of CH₄ for calibration purposes and a blocking filter
190 prevents any remaining 1064 nm radiation from exiting the aircraft. The main OPO output
191 beam is directed through a beam expander to reduce its divergence. The maximum output energy
192 of the OPA transmitter exiting the aircraft was approximately 250 μJ per pulse. The measured
193 linewidth of the OPO was less than 300 MHz but the measurement was limited by the resolution
194 of the Fabry-Perot etalon we used³⁹.

195 The divergence for both laser transmitters was ~150 μrad. Prior to exiting the aircraft through
196 the nadir port, a wedged beam splitter sends a small portion of the outgoing beams (~4 %) to an
197 8.9 cm diameter integrating sphere with two InGaAs detectors attached to one of its ports (one
198 for the OPA and one for the OPO). The detectors measure the outgoing energy monitor pulses

for the OPO or OPA and are digitized by the data acquisition system. The energy monitor pulses are used to normalize the reflected pulses from the ground every 1/16 second in post processing. In addition to the energy monitors, a multimode 200- μm core fiber is also connected to a port of the integrating sphere and collects a small fraction of the outgoing laser energy. The multimode fiber output is collimated and fed back into the receiver telescope and on focused on our sensitive detector to provide a zero range pulse (or “start pulse”) for our ranging algorithm. The time of flight from the zero range pulse to the reflected pulses from the ground is used to determine the IPDA lidar range.

The laser pulses reflected from the ground are collected by a commercial 20 cm diameter receiver telescope (Vixen VC200L) with an effective focal length of 2 m and are coupled into an anti-reflection (AR) coated 600- μm core multi-mode fiber. The receiver field of view (FOV) was 300 μrad . The receiver fiber output is collimated by a lens and directed through a 0.8 nm (FWHM) band pass filter, and then focused onto a HgCdTe enhanced avalanche photodiode (e-ADP) by DRS Technologies⁴³⁻⁴⁵. The detector is a 4x4-pixel array, with the pixel pitch being 80 μm with no gaps between pixels. The detector is operated at 80K and its electrical bandwidth is ~7 MHz.

The signals from the frequency counter, reference cell (OPA or OPO), the energy monitor (OPA or OPO), the zero range pulse, and ground return pulses are digitized by a National Instruments[®] PXI-based data acquisition system containing a FlexRIO FPGA Module, a FlexRIO Digitizer Adapter Module, a Timing and Synchronization Module, and a Global Positioning System (GPS) module. All signals are averaged every 1/16 second and the files are time stamped by the GPS time. Additional averaging can be performed in post-processing. The major parameters of the airborne IPDA lidar are summarized in Table 1 below.

Table 1 Instrument Parameters

Parameter	OPA	OPO
Center Wavelength	1650.958 nm	1650.958 nm
Number of wavelengths used	20	5
Transmitter Energy/pulse	~25-30 μ J	~250 μ J
Transmitter Pulse rate	10 kHz	5 kHz
Transmitter divergence	~150 μ rad	~150 μ rad
Spectral Linewidth	~500 MHz	<300 MHz [*]
Number of seed lasers used	2	5
Pump laser	Burst mode Yb Fiber	Single pulse Nd:YAG
Pump laser energy	350 μ J	1.2 mJ
Receiver diameter	20 cm	20 cm
Receiver Field of view	300 μ rad	300 μ rad
Receiver band pass	0.8 nm (FWHM)	0.8 nm (FWHM)
Detector	4x4 HgCdTe e-ADP	4x4 HgCdTe e-ADP
Detector Pixel Pitch	80 μ m	80 μ m
Detector QE	~ 90%	~ 90%
Detector Temperature	80K	80K
Detector bandwidth	7 MHz	7 MHz
Averaging time	1/16 sec	1/16 sec

^{*} Linewidth measurement limited by the resolution of the scanning Febry-Perot etalon used.

3 Airborne Demonstration Results

3.1 Flights

In late September 2015, the instrument was installed on the NASA DC-8 airborne laboratory, based at Armstrong Flight Research Center Science Aircraft Integration Facility (SAIF) in Palmdale, CA. The transceiver structure supported two small, vibration isolation, optical benches for the OPO and OPA, the receiver telescope, and the transmit optics components. A vibration isolation mechanism for the entire structure minimized the impact of aircraft vibrations. The overall transceiver dimensions were approximately $0.9 \times 2.0 \times 0.8$ m³ and the total weight was 363 kg (Fig. 4). Two aircraft racks on either side of the transceiver structure held ancillary instrumentation needed for the operation of the instrument (data acquisition and control computers, detector, seed lasers, electronics, chillers for the pump lasers, etc.).



Fig. 4. The GSFC IPDA lidar installed on the NASA DC-8 airborne laboratory, in Palmdale, CA. The transceiver structure supported two optical benches for the OPO and OPA, the receiver telescope, and the transmit optics components. The overall transceiver dimensions were approximately $0.9 \times 2.0 \times 0.8 \text{ m}^3$ and the total weight was 363 kg. Two instruments racks on either side of the transceiver contained the control and data acquisition electronics.

A Picarro in-situ analyzer (Picarro G1301-m) measuring methane, carbon dioxide, and water vapor using Wavelength-Scanned Cavity Ring Down Spectroscopy was also installed at a different location in the aircraft to provide in-situ CH_4 reference measurements.

Three flights in the western United States were carried out in late September-early October 2015. Flight planning was constrained by the limited number of flight hours available, the inclement weather and aircraft maintenance issues. Each flight lasted about 4 hours and included several segments at increasing altitudes from 2 to 13 km over varying topography, ground reflectivity (including ocean), and atmospheric conditions. In addition, a spiral descent from ~13 km to near the surface (~30-300 m depending on Federal Aviation Administration flight clearances) was

included in the flight plan in order to sample the localized vertical profile of the CH₄ mixing ratio and associated meteorological parameters (pressure, temperature, humidity, etc.) using the Picarro in-situ sensor and the aircraft's data acquisition system. The IPDA lidar was always turned off below 1 km above ground level (AGL) to comply with strict laser safety requirements. Fig. 5 summarizes our flight paths in the western US. The flight tracks and locations were chosen to minimize the transit flight time and targeted areas of potential CH₄ emission sources.

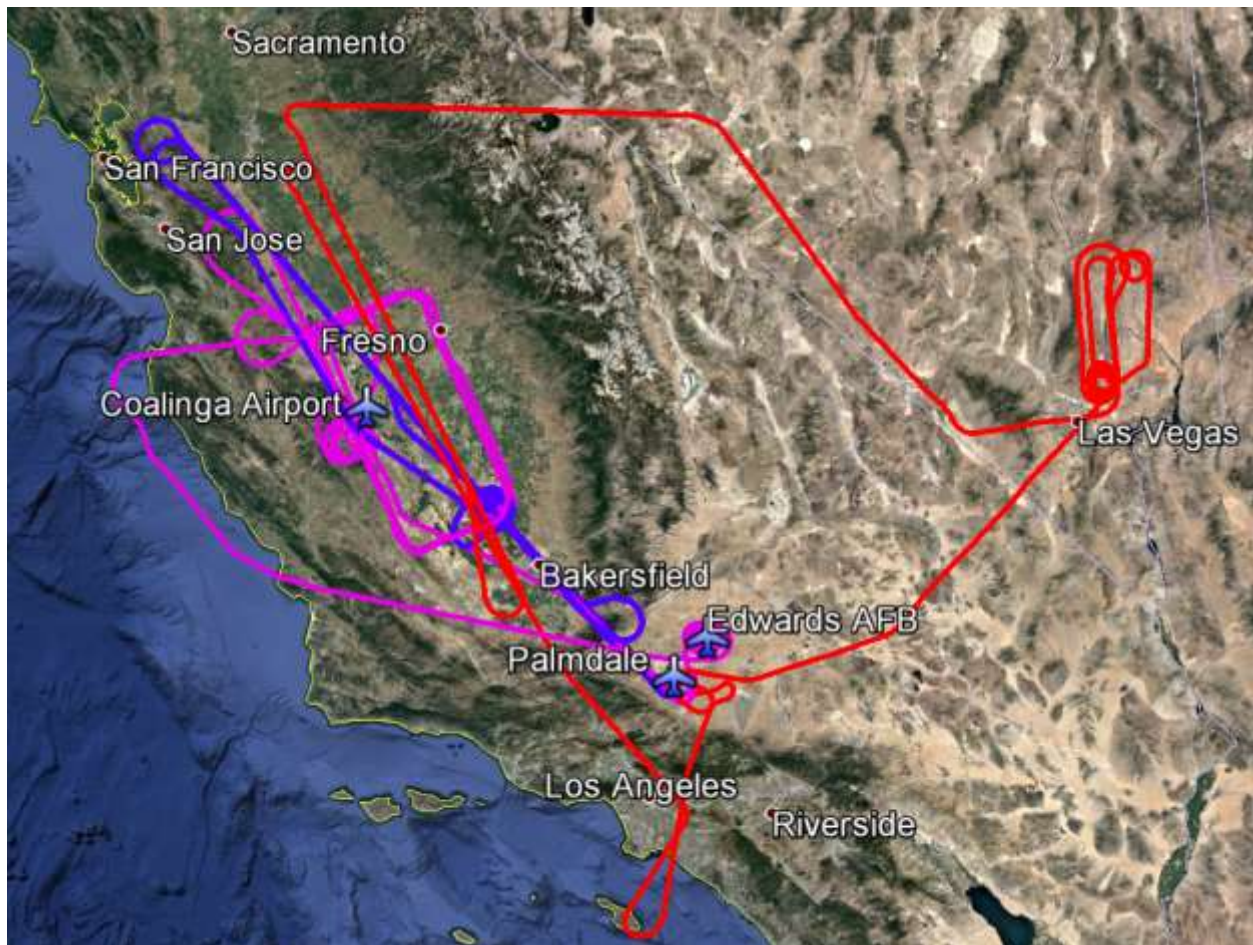


Fig. 5 Flight tracks for the 2015 flights. Flight 1 (blue), Flight 2 (red), Flight 3 (magenta).

For the first two flights, we used the OPA transmitter and for the third flight, we used the OPO. The first flight was mostly over the Central (San Joaquin) Valley of California. We flew on a general south-north track, at three different altitudes at approximately 3.1, 5.9-6.0 km and 12.7

262 km. A large portion of the flight was in and over a dense cloud cover and the spiral descent
263 originally scheduled over Coalinga airport, CA was moved to approximately 40 km north of
264 Bakersfield, CA due to the weather conditions. The Coalinga airport was originally chosen due
265 to its proximity to a large feedlot.

266 The second flight targeted a large landfill approximately 30 km northeast of Las Vegas, NV.
267 After an initial pass at ~10 km and a subsequent spiral descent and low pass over the landfill at
268 ~300 m above ground level (AGL), two more flight segments were flown at 3.2 and 6.4 km.
269 Then we transited over to the Central Valley, CA where we did two high altitude north-south
270 flight segments at 12.7 and 13.1 km. Part of the Central Valley was completely covered by a
271 dense cloud cover during our flight.

272 The third flight was again over the Central Valley of California mostly due to adverse weather
273 conditions and flight restrictions at other candidate sites. We flew on a $\sim 75 \times 160$ km² rectangular
274 path centered on the Central Valley at three different altitudes: 3.1, 6.3 and 12.7 km. A spiral
275 descent and low pass (~30 m AGL) was performed over Coalinga airport. Following the high
276 altitude segment at 12.7 km, we flew over the Pacific Ocean and performed another spiral
277 descent over Edwards Air Force Base (AFB), CA prior to landing in Palmdale.

278 The flight altitude profiles (GPS altitude and IPDA lidar range vs. time) are shown in Fig. 6.

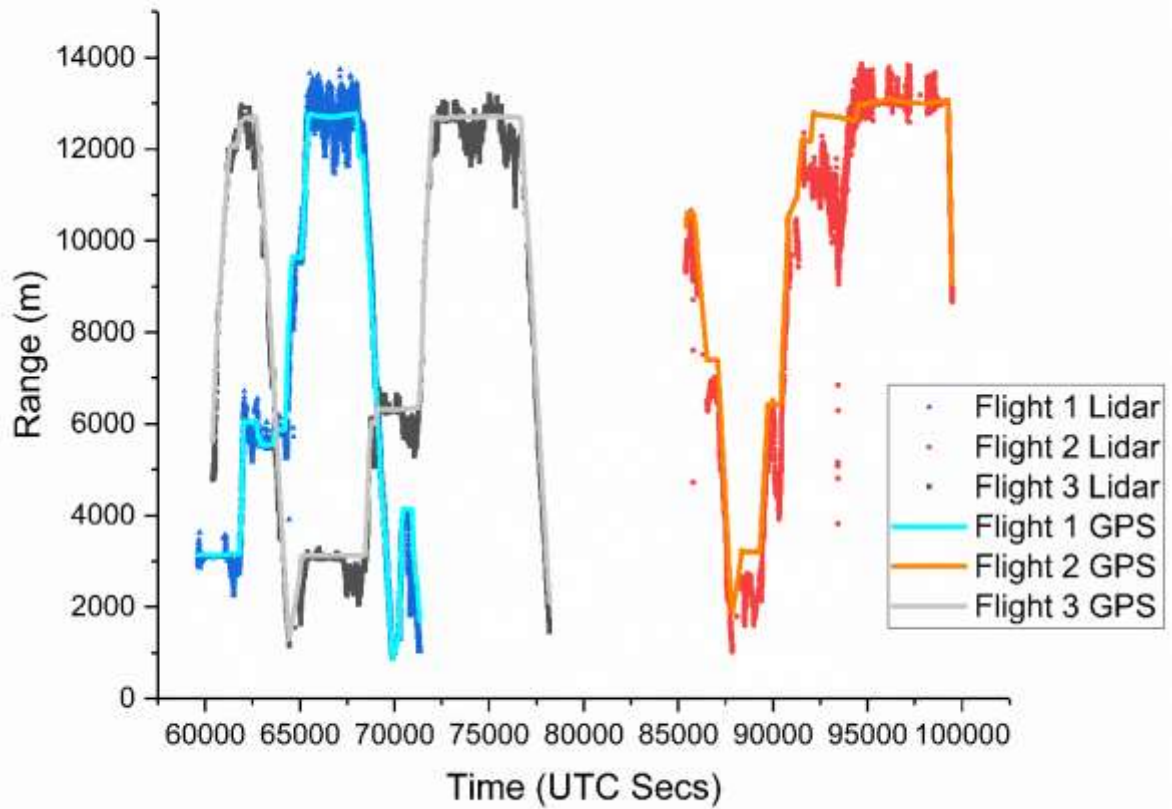


Fig. 6. Flight altitude (range) in m vs. time in UTC seconds since midnight as measured by the GPS receiver and our IPDA lidar for all three flights. The differences between the GPS and the IPDA lidar range are due to topography. The GPS measures the altitude above the mean sea level (or reference ellipsoid) but the IPDA lidar range is the altitude (range) above ground, which includes the topography.

3.2 Retrievals

Our retrieval algorithm uses a least squares fit to minimize the root mean squared error between the IPDA lidar measurements and the model prediction and is similar to the approach used by Abshire et.al.⁴⁶ in their CO₂ retrievals. The averaging time for the data acquisition system is 1/16 sec but the data is further averaged in post processing in 1-sec intervals. First, the range (path length) from the aircraft to the surface is determined from the laser pulse time of flight (TOF) by correlating the first return pulse with the zero range pulse and measuring the time delay of the correlation peak, following the cross-correlation approach by Amediek⁴⁷. The

aircraft is equipped with a GPS antenna and a radar altimeter and we compare our lidar range with the radar range to ensure that that only valid ground return pulses are used in the retrievals. If only cloud returns are present, the data is not used in the analysis. However, there were many occasions we had multiple returns from cumulus and cirrus clouds, and the ground. Segments that did not contain enough valid ground returns due to the presence of clouds or other instrument issues were excluded from the analysis. Generally when fewer than 50% ground returns are present over a 1-second averaging period the data were not used. The algorithm then estimates the column average of CH₄ transmittance of the atmospheric column by fitting the integrated pulse returns from the surface at each wavelength, after normalization by the transmitted pulse energy, the filter transmission, and other instrument calibrations. The algorithm compares the experimental with the theoretically calculated transmittance values and adjusts the fit parameters, including the mixing ratio, to minimize the fit error. The theoretical calculations used a Voigt lineshape, the lineshape parameters from the HITRAN 2008 database and line-by-line radiative transfer calculations⁴⁸. The impact of more complicated lineshape functions and line mixing was not included in the calculations. However, recent spectroscopic measurements by Delahaye⁴⁹ for the MERLIN line at 1645.55 nm suggest that differences of 1.5% up to 5% in the lineshape may arise if these effects are not taken into account. Fig. 7 shows a theoretical CH₄ lineshape from a 400 km altitude orbit with a US standard atmosphere, and a comparison of the corresponding wavelength sampling by the OPA and OPO (20 vs. 5 wavelengths).

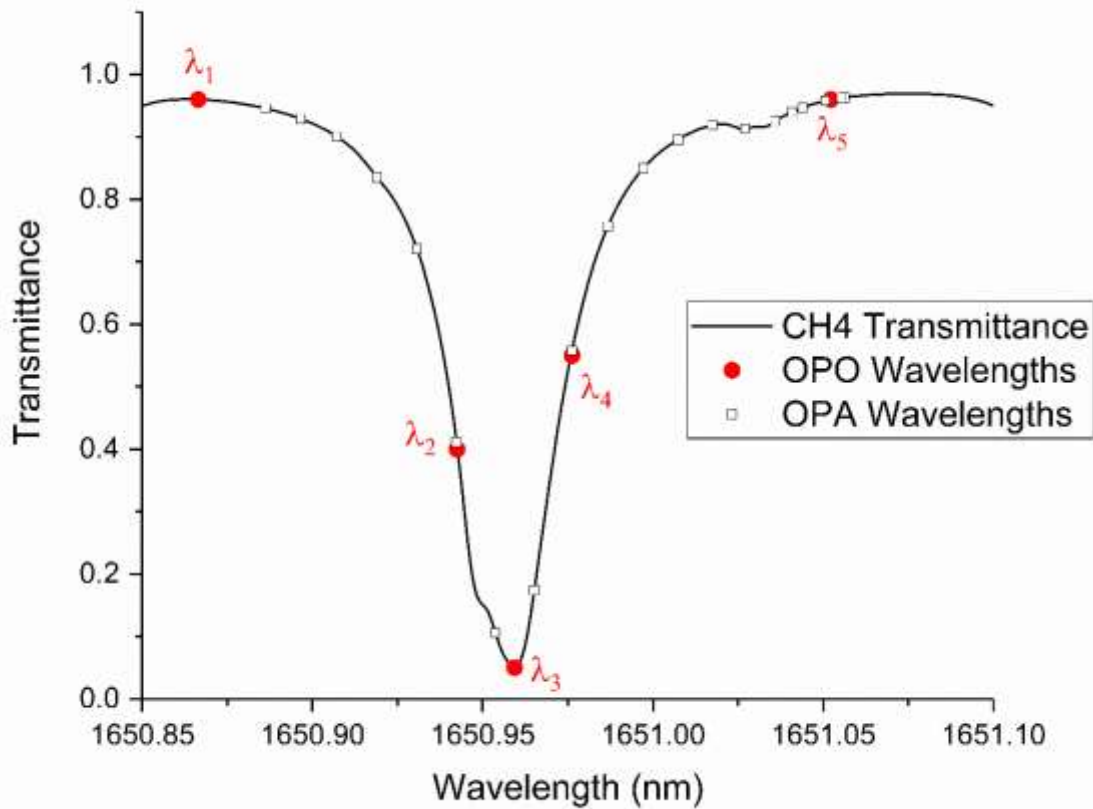


Fig. 7 Theoretical CH₄ transmittance from a 400 km altitude with a US standard atmosphere and a comparison of the approximate wavelength sampling by the OPA (black open squares, 20 wavelengths) and OPO (red solid circles, 5 wavelengths). For clarity, the OPO wavelengths are labeled λ_1 , λ_2 , λ_3 , λ_4 , λ_5 .

The meteorological data for the vertical profile of the atmosphere are obtained from the spiral descents, the Goddard Modeling and Assimilation Office (GMAO) Modern Era Retrospective - Analysis for Research and Applications (MERRA)⁵⁰ and the Goddard Earth Observing System Model, Version 5 (GEOS-5)⁵¹ with a sampling-interpolating interval of 1 second. The true CH₄ mixing ratio profile over the entire flight path is of course, unknown. For simplicity, in our analysis the CH₄ mixing ratio used in the radiative transfer calculations was set at a constant 1900 ppb. Although this value is clearly not an accurate estimate of the true mixing ratio for an entire flight, it is not very different from the column average values obtained by the in-situ spectrometer during our spirals and it provided a reasonable basis for estimating the precision,

but not necessarily the accuracy, of the IPDA lidar. In order to better assess the accuracy of an IPDA lidar more frequent spirals and/or data from radiosondes are needed to infer the mixing ratio profile over a flight path.

Fig. 8 shows the time series data of the in-situ CH_4 mixing ratio measured by the Picarro, the theoretical CH_4 mixing ratio, set at 1900 ppb, and the CH_4 mixing ratio values obtained from our retrieval algorithm for flight 1. There are obvious difference between the instrument retrievals and those from the Picarro. That is to be expected since the Picarro is an in-situ measurement and the lidar measures the column average.

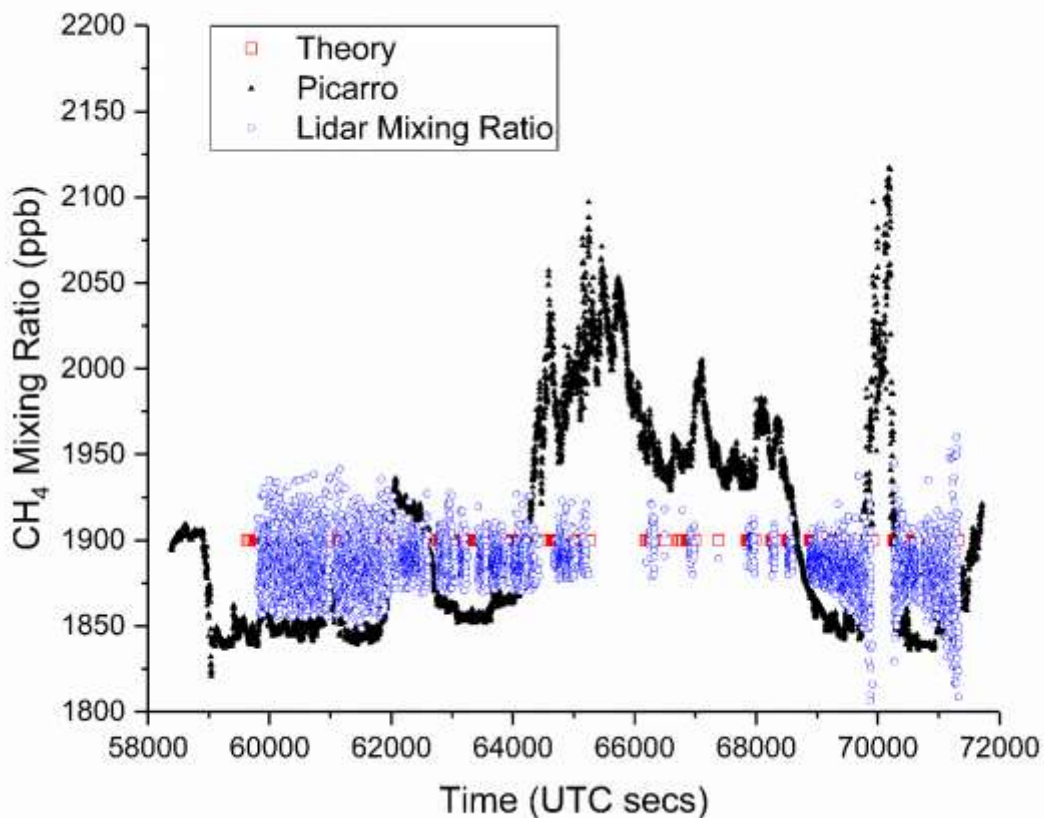


Fig. 8. Time series data for flight 1 showing the in-situ CH_4 mixing ratio as measured by the Picarro, the theoretical CH_4 mixing ratio (set at 1900 ppbv), and the CH_4 mixing ratio values by our retrieval algorithm. Outliers due to instrument adjustments and clouds were filtered out.

339 There were multiple outliers in our retrievals that were filtered out. These were mainly due to
340 two factors: instrument adjustments and optically thick clouds. Due to aircraft vibrations and
341 changes in the cabin temperature and pressure, the OPA transmitter needed alignment
342 adjustments during flight to re-optimize its power. In addition, the wavelength locking circuitry
343 and instrumentation that locked and reported the wavelength of each laser pulse, by measuring
344 the frequency of the beat note between the reference laser and the scanning seed laser
345 wavelengths, would occasionally report erroneous values. The average wavelength values are
346 used if the reported wavelength value did not deviate more than ± 20 pm from the moving
347 average wavelength value. Finally, a significant part of the flight was over broken, optically
348 thick clouds. When all the outliers due to broken clouds and instrument adjustments were
349 filtered out, the agreement between the theoretical CH₄ mixing ratio, set at 1900 ppb, and the
350 retrieved CH₄ mixing ratio values was very good and the standard deviation of the retrieved CH₄
351 mixing ratio was 14.9 ppb or ~0.8% (14.9 ppb/1900 ppb). Assuming the column average CH₄
352 mixing ratio did not vary significantly this value represents a reasonable estimate of the
353 measurement precision of our IPDA lidar. Another good way to assess the IPDA lidar
354 performance is to plot the experimentally retrieved differential optical depth (DOD) vs. the
355 theoretical value. The theoretical and experimental DOD values are determined by the
356 difference in optical depth (OD) between the on wavelength interpolated at the wavelength
357 closest to the peak (~1650.965 nm) and the average value of the OD at the off wavelengths to the
358 left and right of the absorption (~1650.887 and ~1651.056 nm respectively). After removing all
359 the outliers due to laser power adjustments, erroneous wavelength values and broken clouds, the
360 DOD Lidar vs. DOD Theory linear fit (Fig. 9) had a slope of 0.98 and an offset of -0.007. The
361 R² value was 0.994.

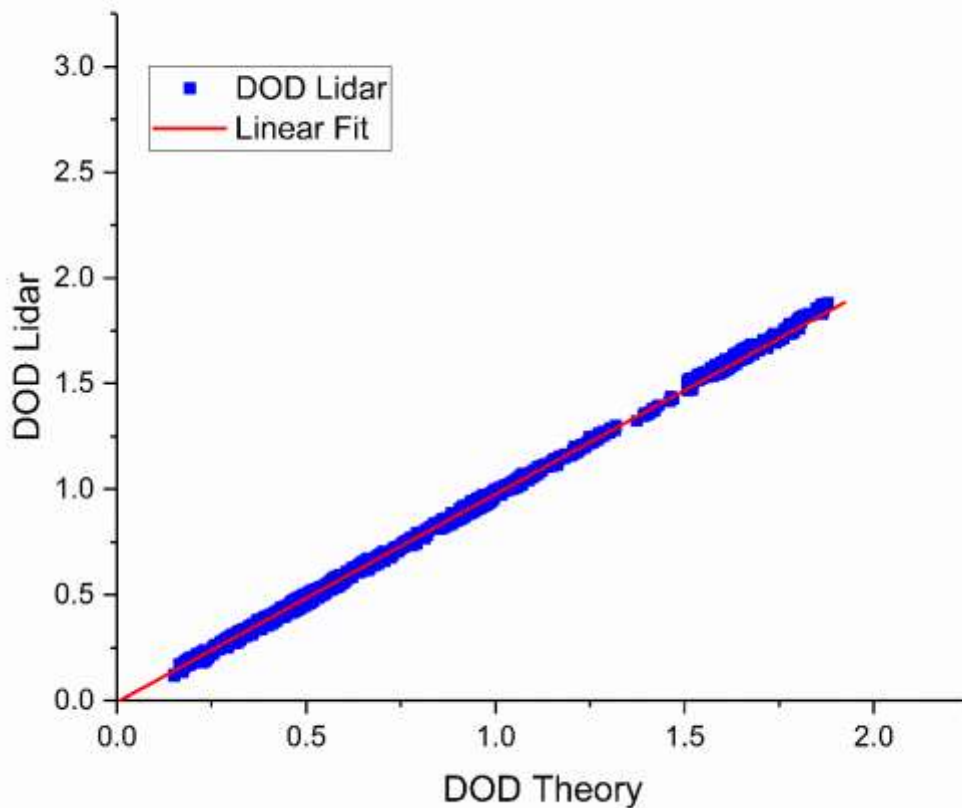


Fig. 9. Lidar DOD vs. Theoretical DOD for flight 1. The linear fit (without the outliers) had a slope of 0.98 and an offset of -0.007 and R^2 value was 0.994.

Fig. 10 Time series for flight 2 data showing the in-situ CH_4 mixing ratio as measured by the Picarro, the theoretical CH_4 mixing ratio (set at 1900 ppbv), and the CH_4 mixing ratio values by our retrieval algorithm. Outliers due to instrument adjustments and clouds were filtered out.

. The standard deviation of the retrieved CH_4 mixing ratio was 13.4 ppb or $\sim 0.7\%$ (13.4

ppb/1900 ppb). The DOD Lidar vs. DOD Theory linear fit for flight 2 had a slope of 0.998 and

an offset of -0.007 (Fig. 11. Lidar DOD vs. Theoretical DOD for flight 2. The linear fit (without

the outliers) had a slope of 0.998 and an offset of -0.007 and R^2 value was 0.990. The R^2 value

was 0.990. These values are consistent with flight 1 results indicating that when the instrument

is operating properly (i.e. no adjustments are being made and no cloud interferences) it is capable

of measuring CH_4 mixing ratios with a 0.7-0.8 % precision. Obviously, the definition of the

instrument “operating properly” is subjective and in our case, it excluded sections of the flights where the instrument needed adjustments. However, the current results provide “proof-of-principle” evidence that a multi-wavelength IPDA lidar with an OPA can provide high enough precision for meaningful science measurements from an airborne platform over a varying topography and altitudes from 2 to 13 km.

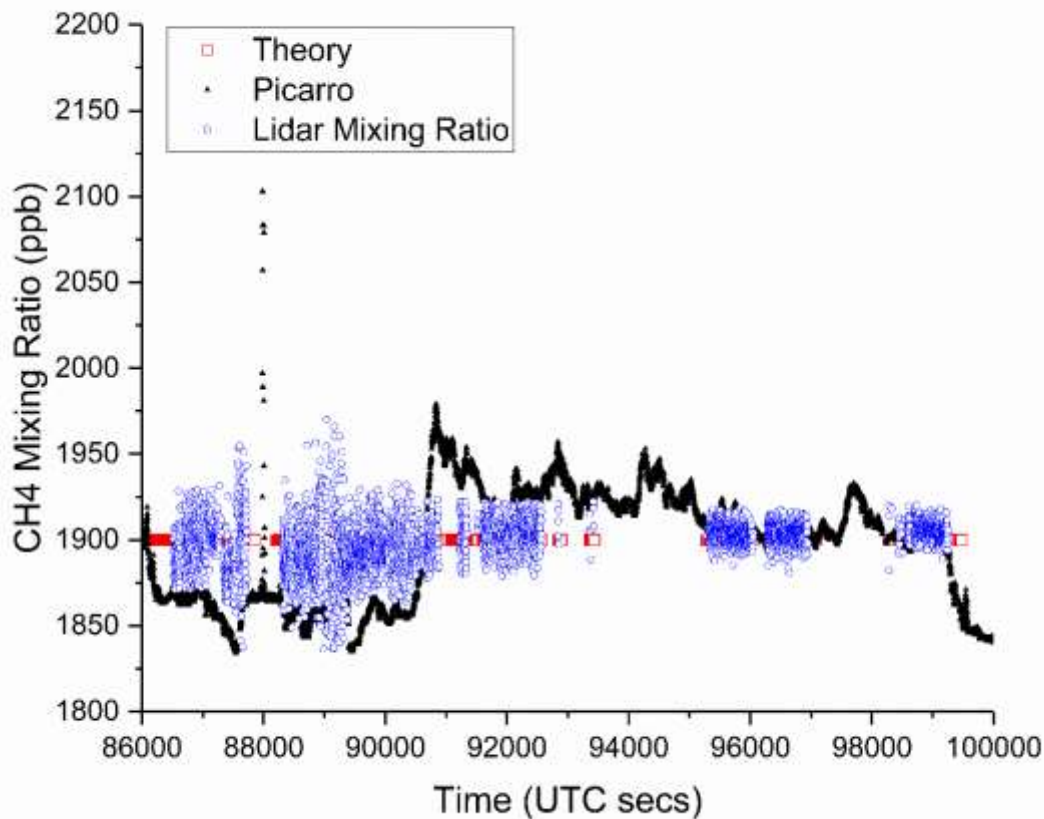


Fig. 10 Time series for flight 2 data showing the in-situ CH₄ mixing ratio as measured by the Picarro, the theoretical CH₄ mixing ratio (set at 1900 ppbv), and the CH₄ mixing ratio values by our retrieval algorithm. Outliers due to instrument adjustments and clouds were filtered out.

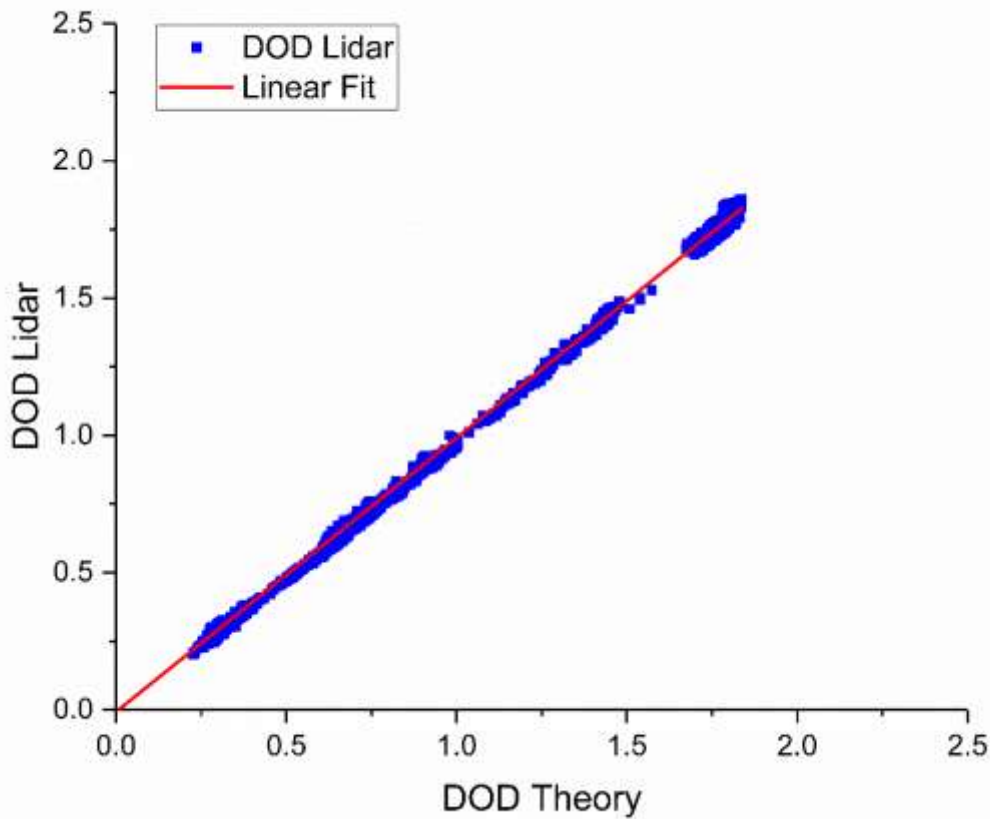


Fig. 11. Lidar DOD vs. Theoretical DOD for flight 2. The linear fit (without the outliers) had a slope of 0.998 and an offset of -0.007 and R^2 value was 0.990.

The last flight (flight 3) used the OPO as the laser transmitter. As shown in Table 1, the OPO used only five wavelengths vs. twenty for the OPA. Fewer wavelengths means that the lineshape is under-sampled and thus, it is more difficult to identify and remove any baseline slope and/or other artifacts in the data. The OPO transmitter also required adjustments during flight. In addition, the high OPO energy ($\sim 250 \mu\text{J}$) presented additional challenges. It saturated our detector especially at lower altitudes. Our initial plan to attenuate the received energy by restricting the receiver aperture size with a variable iris did not work for the flight configuration on the DC-8, even though we tested the idea successfully in the laboratory. The hardware that

was used to restrict the receiver aperture size produced a large near field backscatter when installed in the aircraft saturating and turning off the DRS detector. As a result, it could not be used for flight and thus, the OPO energy was too high for the detector, especially for the low altitude flight segment and over highly reflective surfaces. The detector gain had to be turned down to its minimum value for part of the flight where the detector was presumed to be non-linear. Our initial analysis showed a large discrepancy between the theoretical and experimental DOD values. More importantly, we expected the discrepancy to be worse for the lower altitude segment of the flight, where the detector was saturated and the gain was turned down to a minimum. Contrary to our initial expectations, the DOD discrepancy was worse for the higher altitude segment of the flight, where the detector was not saturated and operating in a linear regime. Repeated post-flight calibrations of the DRS detector in the laboratory failed to uncover any significant detector non-linearities within the digitizer dynamic range (1.23 V peak-peak) that could account for the discrepancy we observed. For a given detector gain (bias) above threshold, the detector is linear over at least two orders of magnitude and even when the detector gain is set to its minimum value the results were repeatable and could be calibrated. Another possible problem we uncovered in our post flight calibration was wavelength locking. Wavelength 1 (λ_1) was initially reporting a “lock” status even though it was not always properly locked on the correct wavelength. The problem was corrected quickly during flight (shortly after the first spiral around 64000 secs UTC) but the data prior to the correction had to be discarded. We hypothesized that the other wavelengths might have also experienced the same issue later in the flight. Several post-flight calibrations in the laboratory with a high-resolution wavemeter showed that the wavelength locking circuitry was operating properly and the circuitry was reporting the wavelength values correctly. A detailed analysis of the OPO reference cell

indicated that the discrepancy was not due to the detector or the wavelength locking. The discrepancy was traced to the fast fiber optic switches used to switch between the five OPO wavelengths. The switches have a small amount of crosstalk and although the crosstalk was initially measured to be relatively small, (~1-1.5%), the effect on the signal can be significant especially at higher altitudes. Because of the crosstalk, the total signal received at each wavelength has contributions from all five wavelengths. Wavelengths 1 and 5 (λ_1 and λ_5) are “off line” and are not absorbed. Wavelengths 2, 3, 4 (λ_2 , λ_3 , λ_4) however, are absorbed and the amount of absorption increases with altitude. As the altitude (absorption) increases the “on line” wavelengths (λ_2 , λ_3 , λ_4) signals, have increasing contributions from the off line wavelengths (λ_1 and λ_5). Thus, a correction factor is needed to account for the crosstalk. The analysis of the reference cell provided the initial evidence and estimate of the crosstalk correction factor. Further refinement of the average correction factor at three different flight altitudes (3.1, 6.3 and 12.7 km) was obtained by ratioing the raw integrated pulse energies at wavelengths 1, 2, 4, and 5 to wavelength 3 and comparing the actual with the theoretical values. Obviously, the correction factor values for different altitudes are just average estimates, not exact values and they vary with altitude and topography. As the aircraft ascends or descends or the topography changes the crosstalk factor will change. Furthermore, as the performance and gain of the OPO cavity changes there is no guarantee that the crosstalk between wavelengths will remain fixed. The crosstalk correction factor accounted for the observed discrepancy for the three constant flight altitude segments of flight 3 (3.1, 6.3 and 12.7 km) and was applied to the analysis. Fig. 12 shows the time series data of the in-situ CH₄ mixing ratio from the Picarro, the theoretical CH₄ mixing ratio, set at 1900 ppb, and the retrieved CH₄ mixing ratio values for the three constant flight altitude segments (3.1, 6.3 and 12.7 km) used in the analysis. When the outliers were

removed, the standard deviation of the retrieved CH₄ mixing ratio was 21.4 ppb or ~1.1 % (21.4 ppb/1900 ppb). The DOD Retrieval vs. DOD Theory linear fit (Fig. 13) had a slope of 1.01 and an offset of -0.003. The R² value was 0.999. These values are comparable but slightly worse to those obtained during flight 1 and 2 with the OPA but the number of outliers that were filtered out was higher.

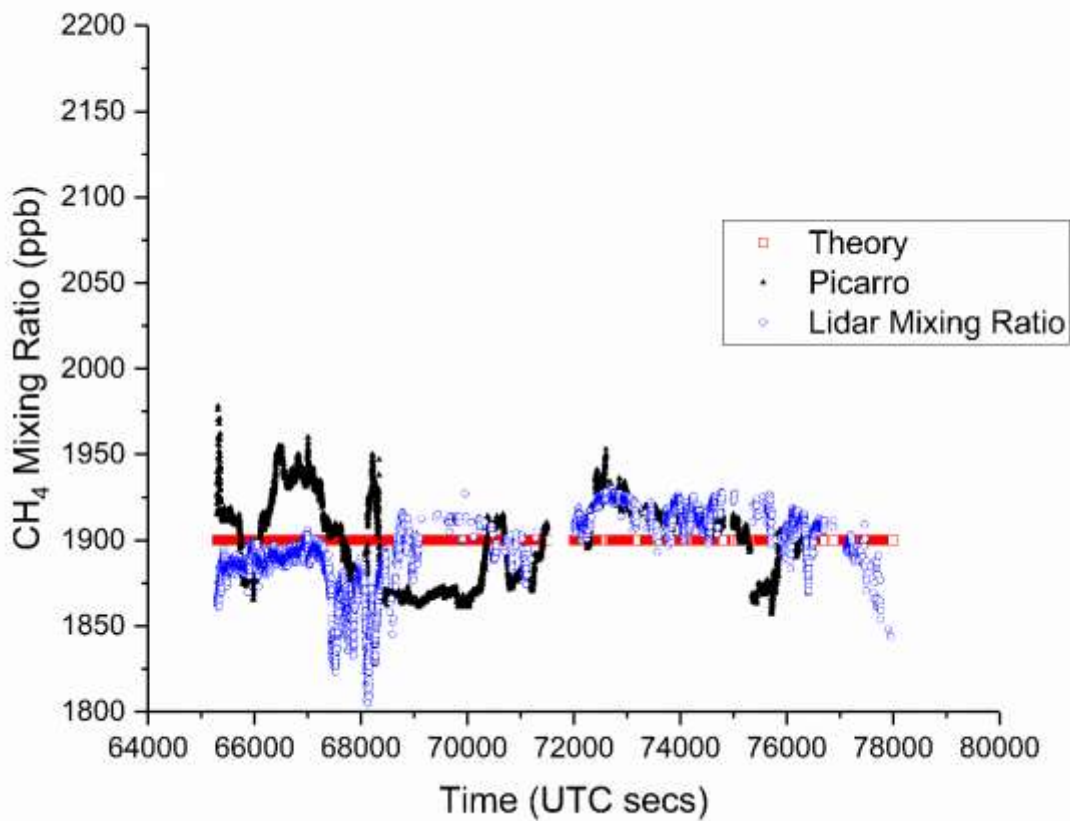


Fig. 12 Time series for flight 3 data showing the in-situ CH₄ mixing ratio as measured by the Picarro, the theoretical CH₄ mixing ratio (set at 1900 ppbv), and the CH₄ mixing ratio values by our retrieval algorithm.

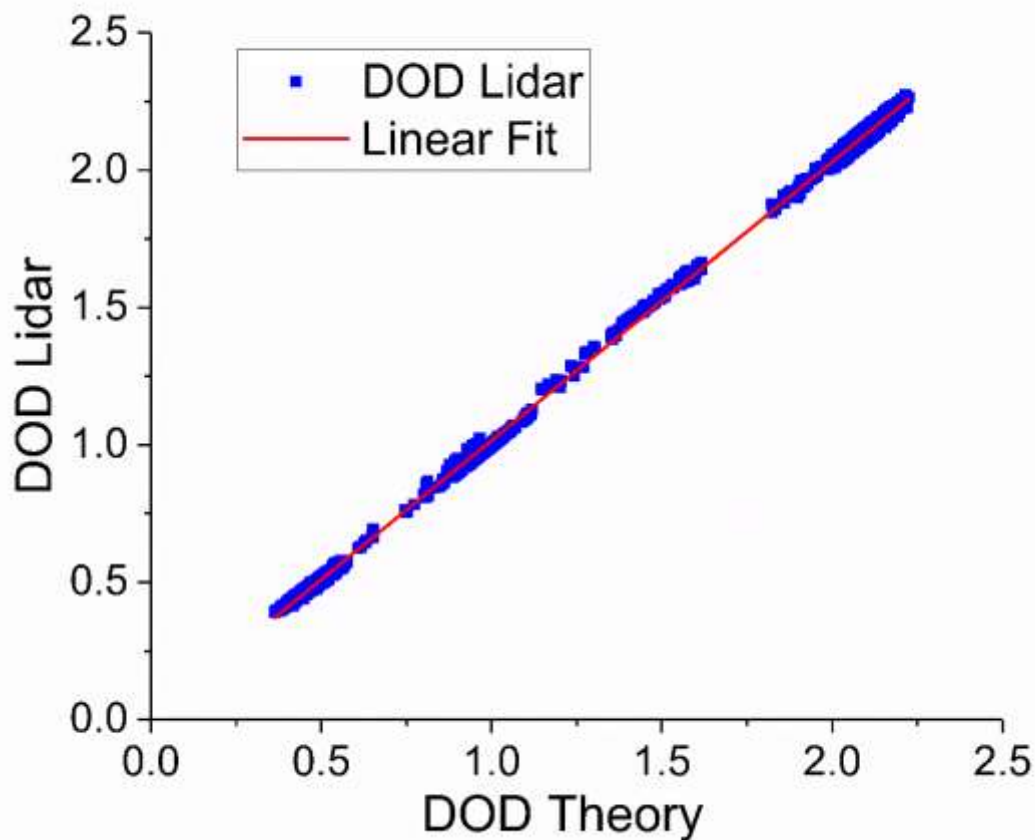


Fig. 13. Lidar DOD vs. Theoretical DOD for flight 3. The linear fit had a slope of 1.01 and an offset of -0.003. The R^2 value was 0.999.

4 Discussion

The high accuracy and precision needed for CH_4 measurements poses several challenges for the IPDA lidar design. Currently the laser transmitter poses the greatest challenge. The transmitter must have narrow linewidth (<100 MHz), must be tunable to scan over the CH_4 absorption line (~ 250 pm), and must have high pulse energy. The exact energy requirement also depends on the detector quantum efficiency, receiver aperture size, and other instrument parameters. Our link margin calculations show that approximately $600 \mu\text{J}$ is needed for space to achieve a 0.5% random error with a 1 m diameter telescope.

467 We have tried to address several potential error sources in our instrument design: Errors due to
468 cloud and aerosol scattering are minimized by our pulsed approach, which digitizes the entire
469 atmospheric column return and gates the returns from the ground. The IPDA lidar cannot
470 penetrate optically thick clouds. However, our ranging algorithm provides accurate knowledge
471 of the total pathlength minimizing the effect of multiple scattering and excluding returns from
472 clouds. In order to improve the accuracy of the measurement better knowledge of the
473 spectroscopic parameters of the CH₄ line and a more sophisticated lineshape function are needed.
474 The CH₄ line we used is actually comprised of multiple lines with different linestrengths and
475 temperature dependence.

476 There are obvious difference between the instrument retrievals and those from the Picarro. The
477 lidar values show less variation in the CH₄ mixing ratio, which is to be expected since the Picarro
478 is an in-situ measurement and the lidar measures the column average. The Picarro recorded a
479 significant increase in the in-situ CH₄ mixing ratio only for flight 1 despite the fact that we tried
480 to target landfills and other areas of increased CH₄ mixing ratios. No significant CH₄ mixing
481 ratio increase was observed near these areas by the Picarro for flights 2 and 3. The biggest
482 increase was observed during the spiral for flight 1 near the ground (around 7000 secs UTC in
483 Fig. 8) when the lidar was turned off to satisfy the laser safety requirements. The other big
484 increase was observed earlier in the flight (around 6600-6700 secs UTC in Fig. 8). The source of
485 that increase is unclear.

486 The predicted values (i.e. “truth”) to which our experimental data are fitted to, assumed a
487 constant CH₄ mixing ratio of 1900 ppb. The actual CH₄ mixing ratio profile over the entire flight
488 path is of course, unknown. Although 1900 ppb is clearly not an accurate estimate of the true
489 mixing ratio for an entire flight, it provided a reasonable basis for estimating the precision, but

not necessarily the absolute accuracy, of the IPDA lidar. In order to assess the accuracy of any IPDA lidar multiple spirals and/or data from other sources such as radiosondes are needed to measure the actual vertical distribution of CH₄ mixing ratio. Even then, differences and biases will remain. Differences in the spectroscopic database (HITRAN 2012 vs. 2008), and effects like line mixing and speed dependent profiles produce different results. Our incomplete knowledge of the state of the atmosphere (pressure, temperature, and humidity) and the accuracy of our meteorological model also contribute. Although the MERRA model was adequate for these initial demonstration flights, in order to evaluate the IPDA lidar accuracy, better knowledge and modeling of the state of the atmosphere and CH₄ vertical profiles at the local level are needed. These can be obtained in future flights by increasing the frequency and location of the spirals, including data from other instruments and radiosonde data if available. Finally, lidar issues due to a variety of factors such as shot noise, ground reflectivity, speckle noise, wavelength and power stability of the laser transmitter, etalon fringes, and in the case of the OPO cross talk between wavelengths all contribute to biases in the measurement. With the limited data obtained from these demonstration flights, we are not able to separate individual bias sources. However, we did observe that the random noise was reduced by the expected $1/\sqrt{t_{av}}$, where t_{av} is the averaging time, for up to ~5-10 secs. After ~5-10 secs, no improvement in the noise statistics is observed.

The CH₄ IPDA lidar needs significant engineering improvements to increase its reliability. The opto-mechanical design, laser transmitter stability, for both the OPA and OPO needs to be considerably improved to reduce the effects of vibration, temperature and pressure. The locking electronics and diagnostics for the seed lasers and OPO cavity also need to improve. The isolation between OPO wavelengths needs to increase by at least an order of magnitude to

eliminate the need for crosstalk correction factors. All of these improvements are feasible if proper engineering resources can be applied.

The OPA, which used 20 wavelengths, produced better fits, was simpler to implement than the OPO because it did not require an optical resonator cavity, and was easier to align and tune. However, it is extremely difficult to scale the OPA energy to that needed for space ($\sim 600 \mu\text{J}$ depending on the receiver size and other instrument parameters) and maintain a narrow linewidth. The highest energy we obtained in the laboratory with our OPA was $290 \mu\text{J}$ using a two-stage OPA, and the burst-mode Yb fiber laser amplified by a custom solid-state amplifier as a pump. However, at high energies, the OPA output spectrum typically consists of sharp peak near the seed wavelength and a broad side lobe, when the parametric gain is high. In that case, we cannot clearly define the linewidth but it is generally too wide for accurate CH_4 IPDA lidar measurements. In addition, for a space mission we are aiming for a simple and efficient single stage - not a complex multi-stage - OPA based on quasi-phase matching (QPM). In this configuration, we have observed that the OPA output linewidth does not fully converge to the seed linewidth, giving wide side lobes, especially when pump and seed fluences are high and low, respectively⁵². Similar side lobes were observed in seeded QPM-based OPA system for a CO_2 lidar⁵³. Back-conversion and parametric amplification of the seed's side lobes are possible causes. Complex OPA/OPG systems in other wavelength regions have been developed with a narrow linewidth⁵⁴. However, it is difficult to predict how they can be implemented with a multi-wavelength IPDA lidar for space because of their complexity. If the seed laser power can be significantly scaled up then it may be possible to achieve energies of $600 \mu\text{J}$ out of the OPA with a narrow linewidth. With the existing seed and pump laser technology, we do not see a path

535 to space for the OPA in the near future. However, it remains a viable transmitter for CH₄
536 measurements from an airborne platform.

537 In the OPO, narrow linewidth was achieved by using an optical resonator cavity, which also
538 enhances the energy of the non-linear conversion. Our 5-wavelength OPO uses a 1.2 mJ GSFC-
539 built solid-state pump laser and a triangular optical ring cavity. We have since replaced the
540 GSFC-built solid-state pump laser with a smaller, compact Yb fiber laser and redesigned the
541 OPO cavity to improve stability. In the laboratory we demonstrated energies of ~250 μ J at 5
542 kHz with a narrow (transform limited) linewidth. The OPO energy could be scaled to space but
543 it requires complex optical phase-lock loops and cavity control.

544 In recent years, resonantly pumped Erbium (Er) doped YAG, Er:YAG and Er:YGG, lasers,
545 which directly emit at 1645.5 and 1650.96 nm, respectively offer another option for a CH₄
546 transmitter. Using Er:YAG for CH₄ detection dates back to 1972⁵⁵ and recent successful
547 demonstration and commercialization of high power and high spectral brightness pump sources
548 have afforded the realization of resonant pumping of Er:YAG⁵⁶⁻⁵⁹ and Er:YGG^{60, 61}. The
549 emission cross-section of Er:YAG crystal is centered near 1645.3 nm and falls off rapidly at
550 1650.96 nm. It is near the MERLIN lines at 1645.55 nm which are relatively wide (~56 pm).
551 Our CH₄ line at 1650.96 nm is narrower (~36 pm) which makes fast tuning easier. Unfortunately,
552 Er:YAG cannot be used as a gain medium at 1650.96 nm but Er:YGG can be used as a potential
553 medium for lasing at that wavelength. Both materials are good candidates for a CH₄ laser
554 transmitter. Power scaling for both materials, multi-wavelength operation and tuning
555 considerations remain.

556 Our preliminary radiative transfer calculations show that both lines (Er:YAG at 1645.55 nm and
557 Er:YGG at 1650.96 nm) have similar temperature sensitivity and are well suited for space born

CH₄ measurements. Recent high accuracy spectroscopic measurements indicate that line mixing effects in the Er:YAG 1645.55 nm line⁴¹ should also be taken into account. We expect similar effects to be present for the Er:YGG line at 1650.96 nm.

5 Summary

We reported on an airborne demonstration of atmospheric CH₄ measurements with an Integrated Path Differential Absorption (IPDA) lidar using an optical parametric oscillator (OPO) and optical parametric amplifier (OPA) laser transmitter and sensitive avalanche photodiode detector. The lidar measured the atmospheric CH₄ absorption at multiple, discrete wavelengths near 1650.96 nm. The instrument was deployed in 2015 aboard NASA's DC-8 airborne laboratory and measured CH₄ mixing ratios from 2 km to 13 km. Relatively high precision measurements of 0.7% to 1.1 % were demonstrated for all three flights however, many areas of improvement remain. The stability and reliability of the laser transmitters need to improve considerably but the basic measurement approach has been demonstrated. We are currently improving our airborne instrument with better opto-mechanical design and compact, more stable laser transmitters. We hope to fly again in the near future when the next opportunity arise.

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Figure Captions

Fig. 1. Two-way atmospheric transmittance near 1650.96 nm a from a 400 km orbit using the 2008 HITRAN database and a US standard atmosphere. The CH₄ line is mostly isolated from adjacent CO₂ and H₂O vapor lines.

Fig. 2. Simplified functional block diagram of our IPDA lidar. The lidar can use one of two different laser transmitters using a movable selection mirror: An Optical Parametric Amplifier (OPA) or an Optical Parametric Oscillator (OPO). The transmitters use DFB diode lasers for seed lasers but different pump lasers. Only one laser transmitter is operating during flight.

Fig. 3. Example of temporal shape of the OPA burst pulse from the energy monitor detector showing individual pulses within the burst pulse.

Fig. 4. The GSFC IPDA lidar installed on the NASA DC-8 airborne laboratory, in Palmdale, CA. The transceiver structure supported two optical benches for the OPO and OPA, the receiver telescope, and the transmit optics components. The overall transceiver dimensions were approximately 0.9×2.0×0.8 m³ and the total weight was 363 kg. Two instruments racks on either side of the transceiver contained the control and data acquisition electronics.

Fig. 5 Flight tracks for the 2015 flights. Flight 1 (blue), Flight 2 (red), Flight 3 (magenta).

Fig. 6. Flight altitude (range) in m vs. time in UTC seconds since midnight as measured by the GPS receiver and our IPDA lidar for all three flights. The differences between the GPS and the IPDA lidar range are due to topography. The GPS measures the altitude above the mean sea level (or reference ellipsoid) but the IPDA lidar range is the altitude (range) above ground, which includes the topography.

Fig. 7 Theoretical CH₄ transmittance from a 400 km altitude with a US standard atmosphere and a comparison of the approximate wavelength sampling by the OPA (black open squares, 20 wavelengths) and OPO (red solid circles, 5 wavelengths). For clarity, the OPO wavelengths are labeled $\lambda_1, \lambda_2, \lambda_3, \lambda_4, \lambda_5$.

Fig. 8. Time series data for flight 1 showing the in-situ CH₄ mixing ratio as measured by the Picarro, the theoretical CH₄ mixing ratio (set at 1900 ppbv), and the CH₄ mixing ratio values by our retrieval algorithm. Outliers due to instrument adjustments and clouds were filtered out.

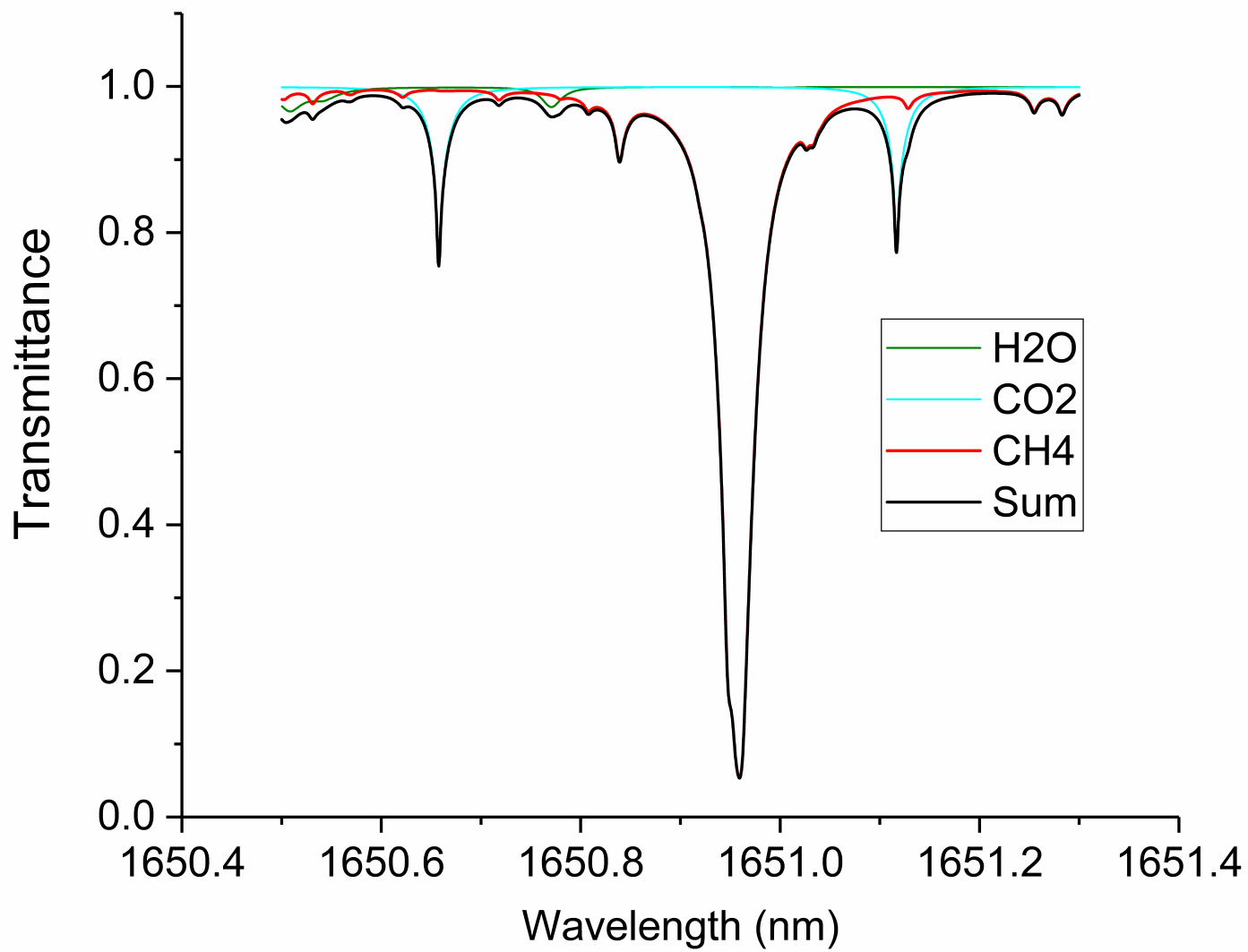
Fig. 9. Lidar DOD vs. Theoretical DOD for flight 1. The linear fit (without the outliers) had a slope of 0.98 and an offset of -0.007 and R² value was 0.994.

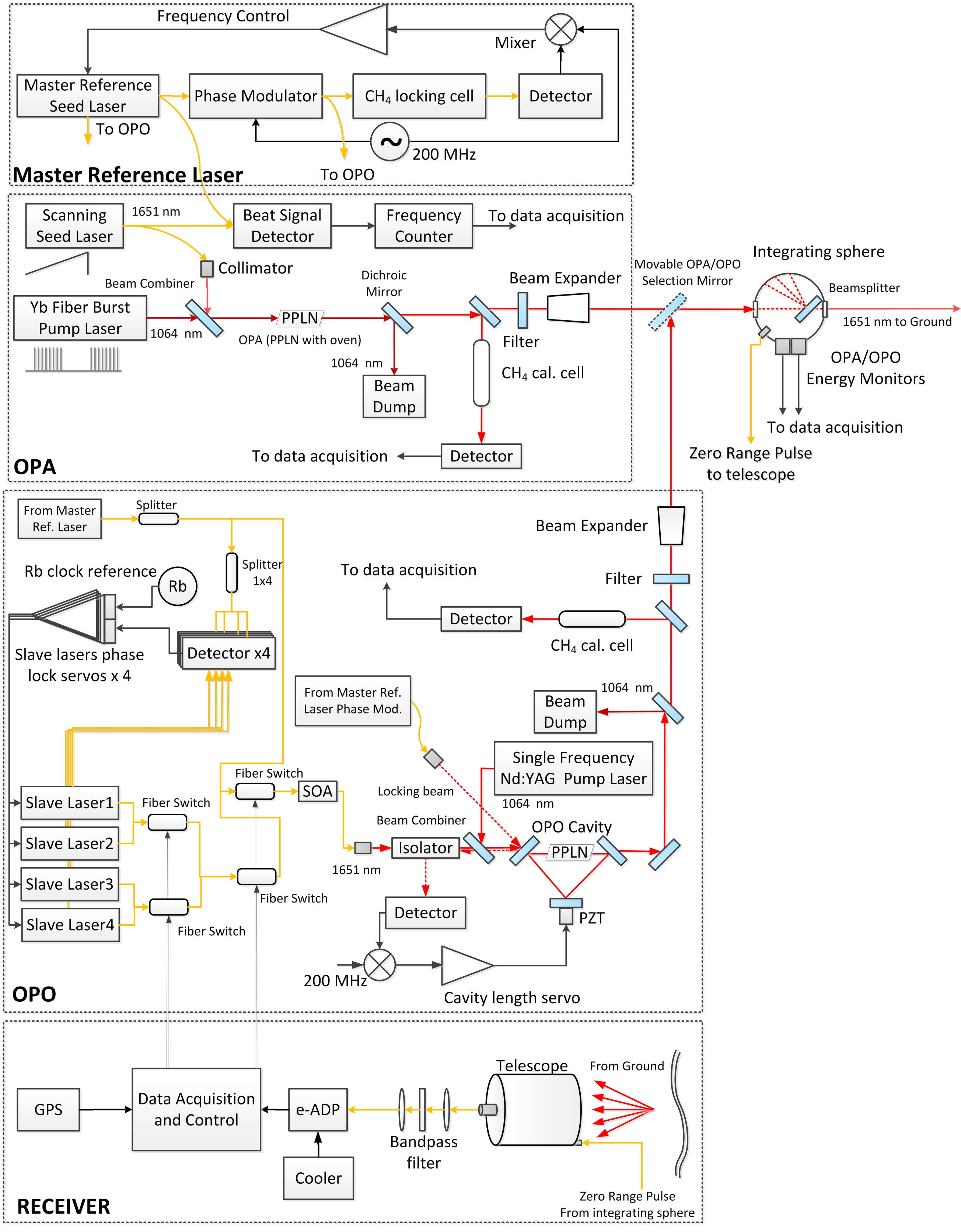
Fig. 10 Time series for flight 2 data showing the in-situ CH₄ mixing ratio as measured by the Picarro, the theoretical CH₄ mixing ratio (set at 1900 ppbv), and the CH₄ mixing ratio values by our retrieval algorithm. Outliers due to instrument adjustments and clouds were filtered out.

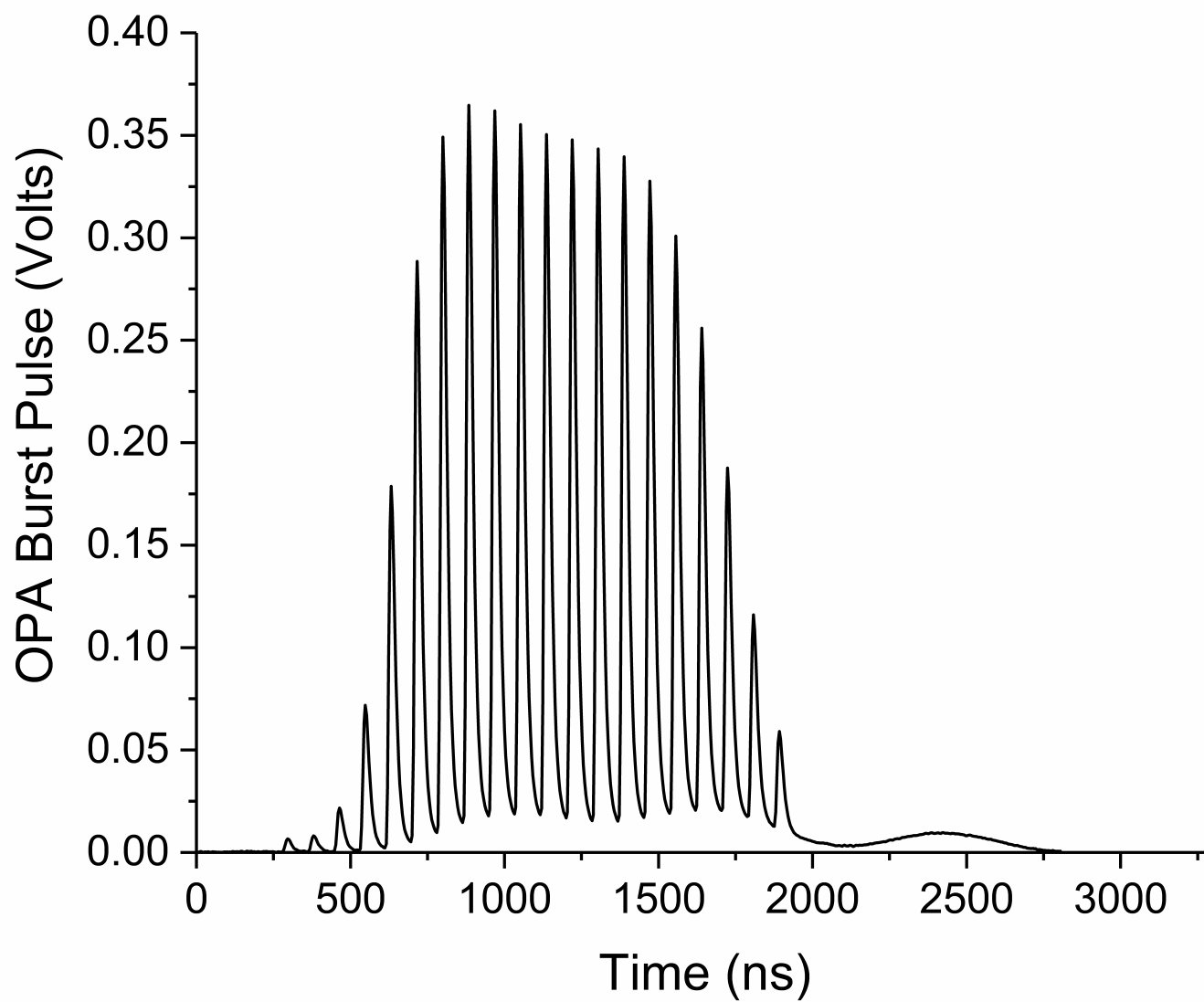
Fig. 11. Lidar DOD vs. Theoretical DOD for flight 2. The linear fit (without the outliers) had a slope of 0.998 and an offset of -0.007 and R² value was 0.990.

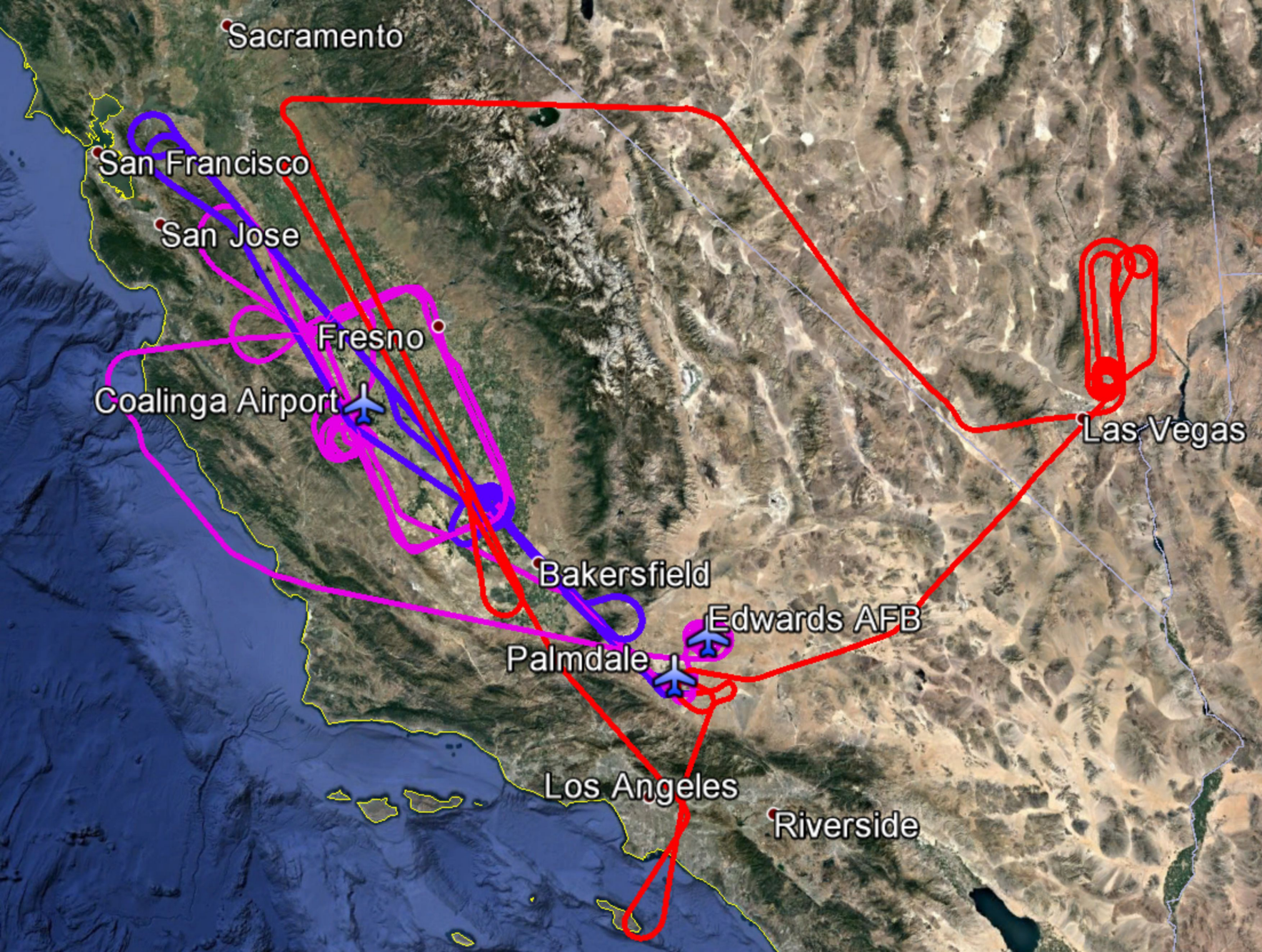
Fig. 12 Time series for flight 3 data showing the in-situ CH₄ mixing ratio as measured by the Picarro, the theoretical CH₄ mixing ratio (set at 1900 ppbv), and the CH₄ mixing ratio values by our retrieval algorithm.

Fig. 13. Lidar DOD vs. Theoretical DOD for flight 3. The linear fit had a slope of 1.01 and an offset of -0.003. The R² value was 0.999.









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